



GEOLOGY OF WYOMING (/)

WIND RIVER CANYON PART TWO

Wow Factor *(4 out of 5 stars):*



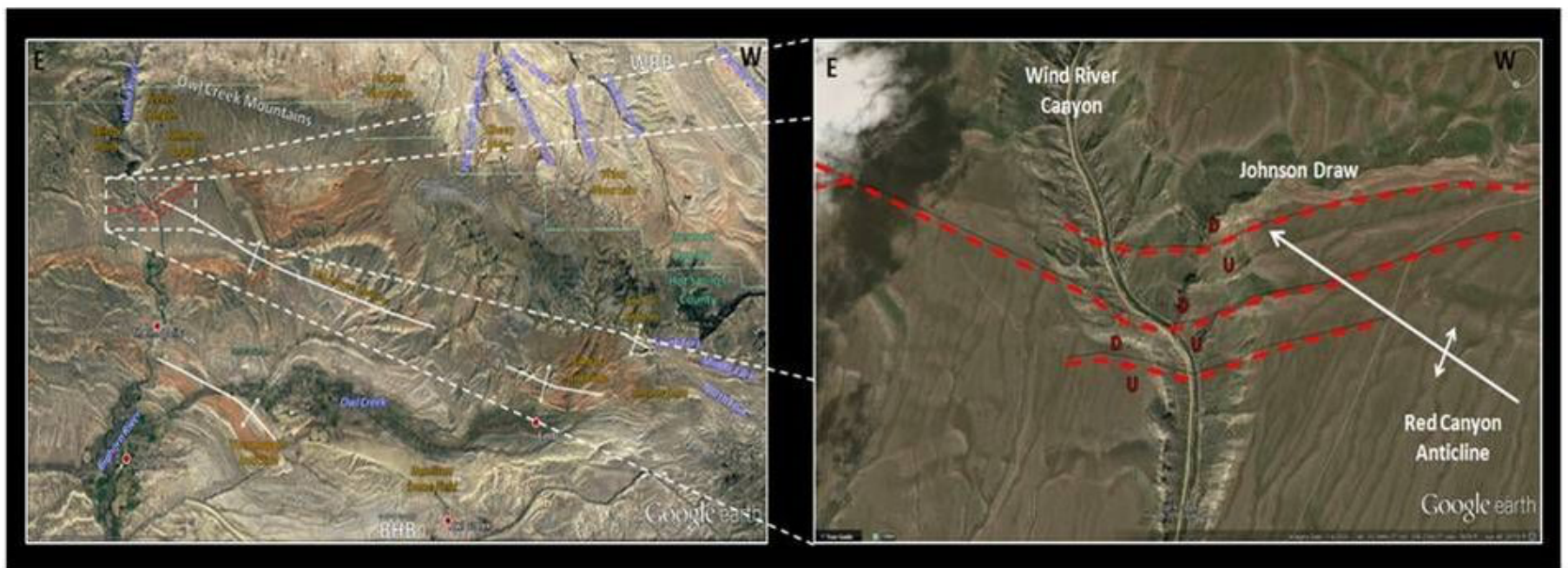
Geologist Factor *(4 out of 5 stars):*



Continued from Wind River Canyon Part One (<http://www.geowyo.com/wind-river-canyon-part-1.html>)

0.2 6.0 0.4 11.8

Top Madison to west (R/L). The Madison has produced oil where the thickness between productive Phosphoria and the Madison is exceeded by the amount of structural closure. Just ahead on the west (R/L) is Johnson Draw. This is the continuation of deformation related to the Red Canyon anticline west of the canyon's northern entrance. A zone of down to the south high-angle normal faults offset the Madison Formation. The vertical displacement is about 50 feet.



Johnson Draw area, Wind River Canyon.

Images: Google Earth

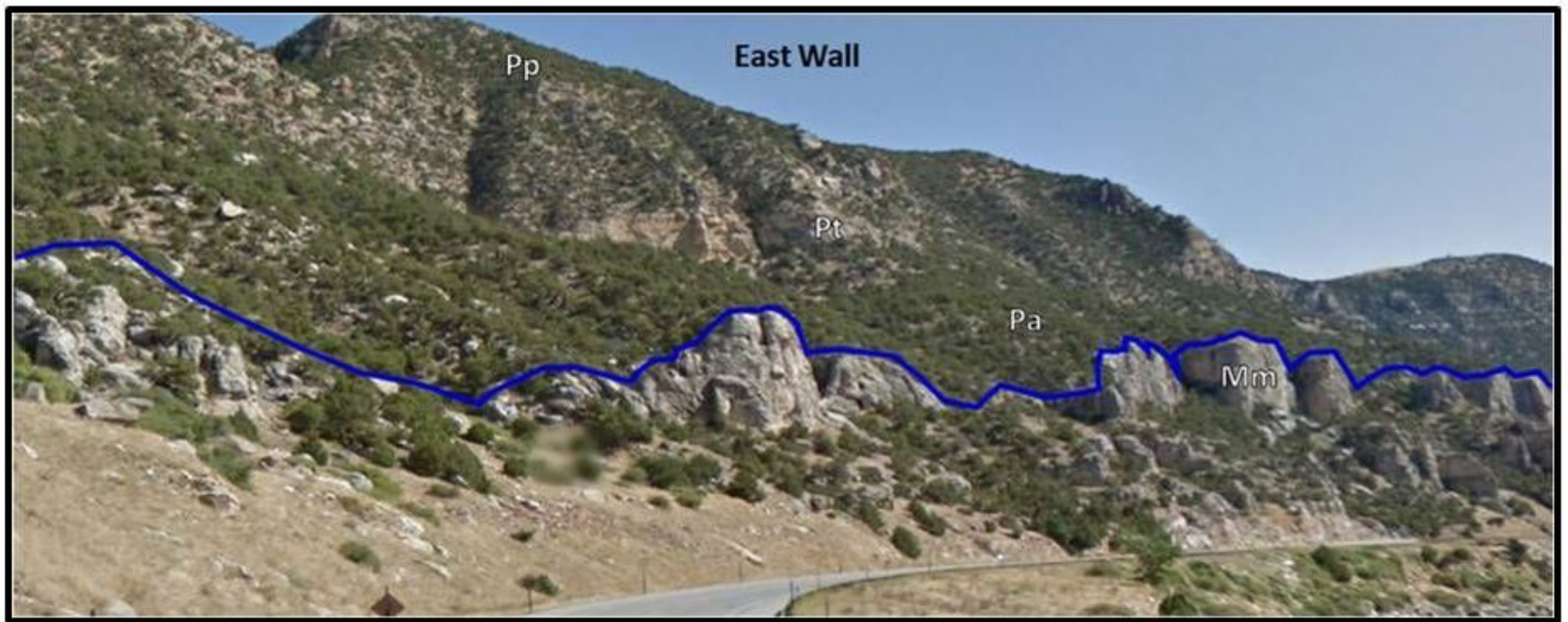
0.4

6.4

1.0

11.4

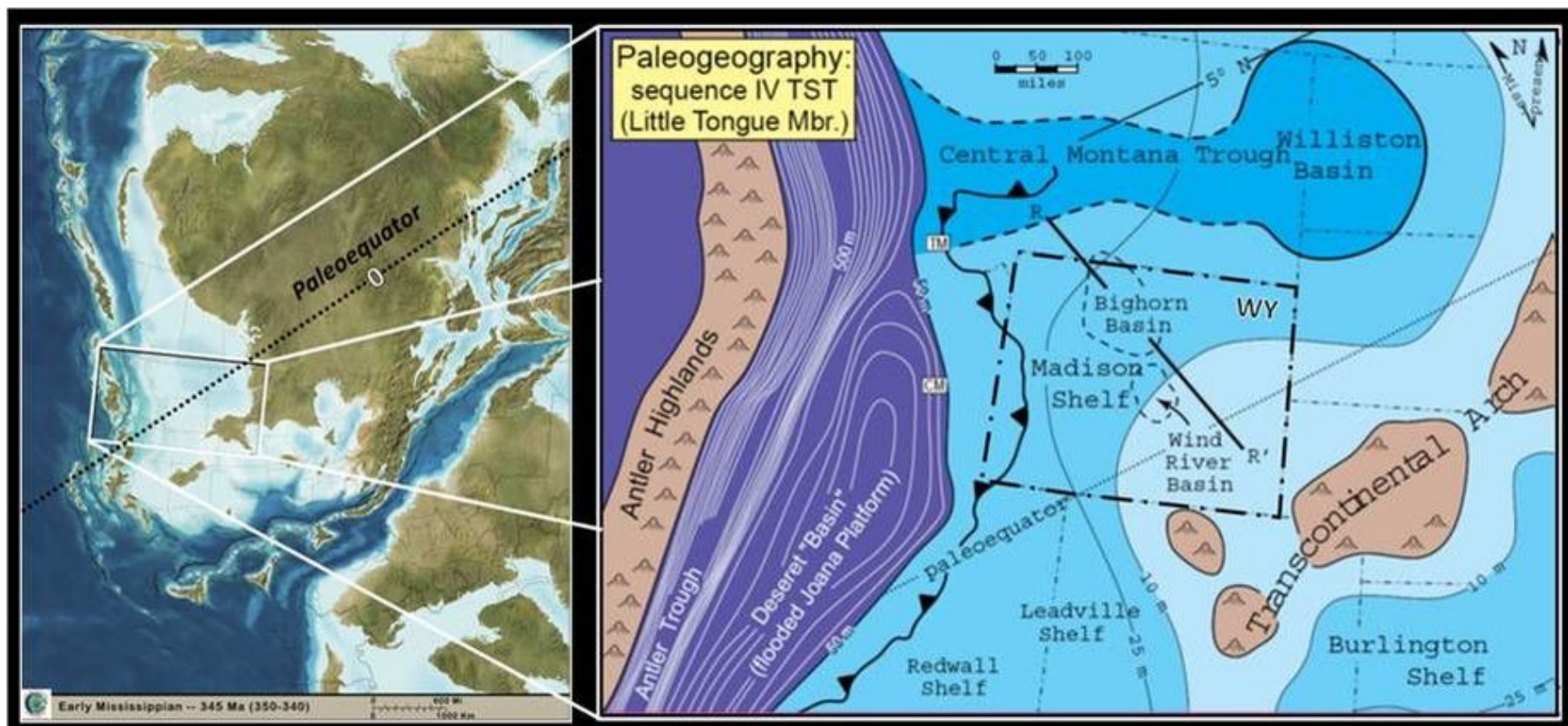
Road veers to the left (east) at Johnson Draw. Excellent view of the Madison erosion surface on the east wall of the canyon where highway turns back south ahead. The Mississippian Madison Limestone is about 475 feet thick. Most of the basal 300 feet is dolomite and the remaining 75 feet is largely limestone. The upper portion of the Madison Limestone is often vuggy and cavernous due to a karst topography which developed during subaerial exposure.



View of erosion surface at top of the Madison Formation near Johnson Draw. Geologic notation: Pp: Permian Phosphoria Formation Pt: Pennsylvanian Tensleep Formation Pa: Pennsylvanian Amsden Formation Mm: Mississippian Madison Formation. This view is at 6.6 miles.

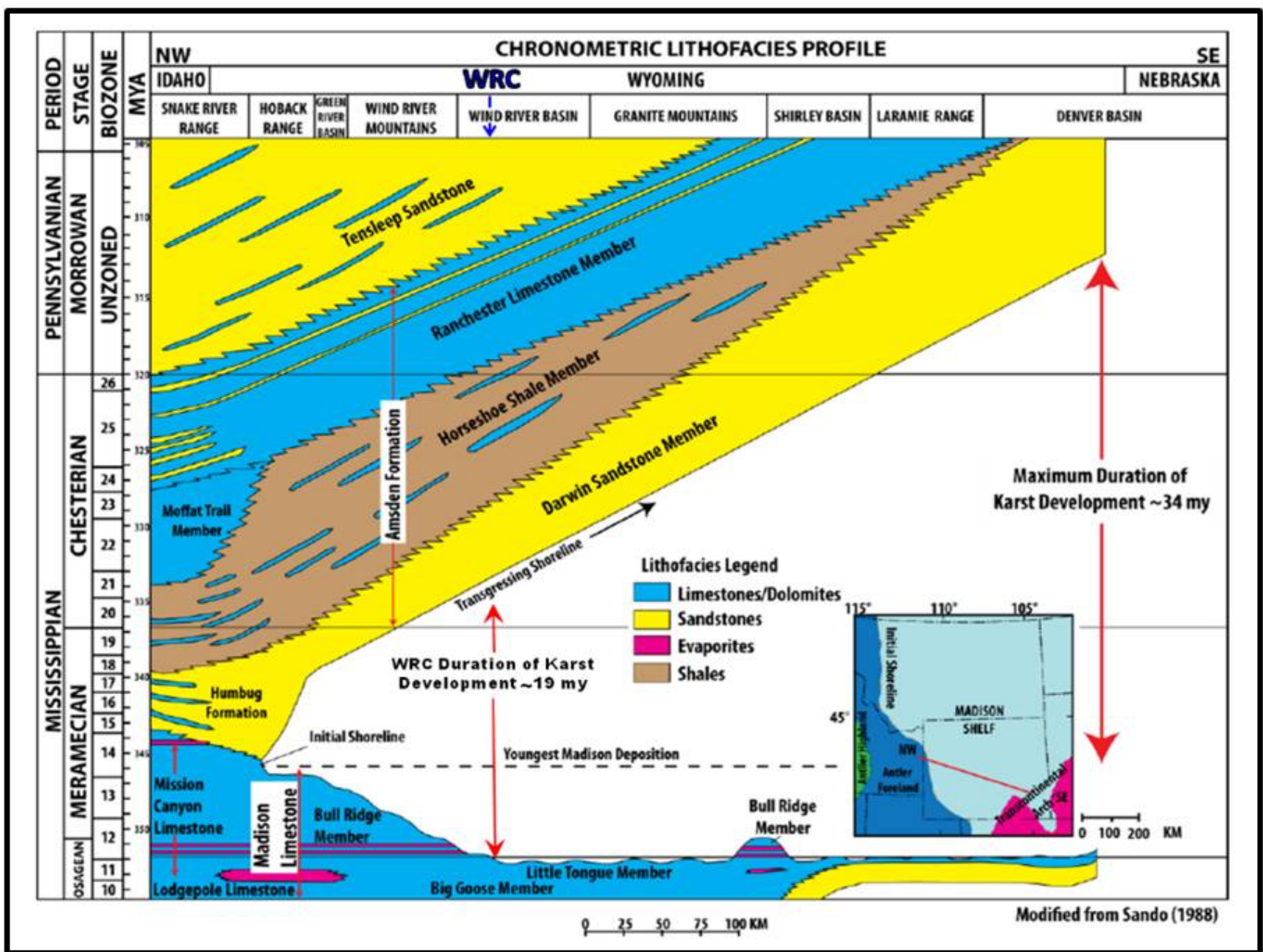
Image: Google Earth Street View

The Madison Formation was deposited on a shallow marine water platform that was part of a large carbonate province located between zero and five degrees north paleolatitude. The platform was bounded on the north by the Central Montana Trough, by the intercratonic Williston Basin to the northeast, the Transcontinental Arch to the southeast, and the Antler Highlands to the west.



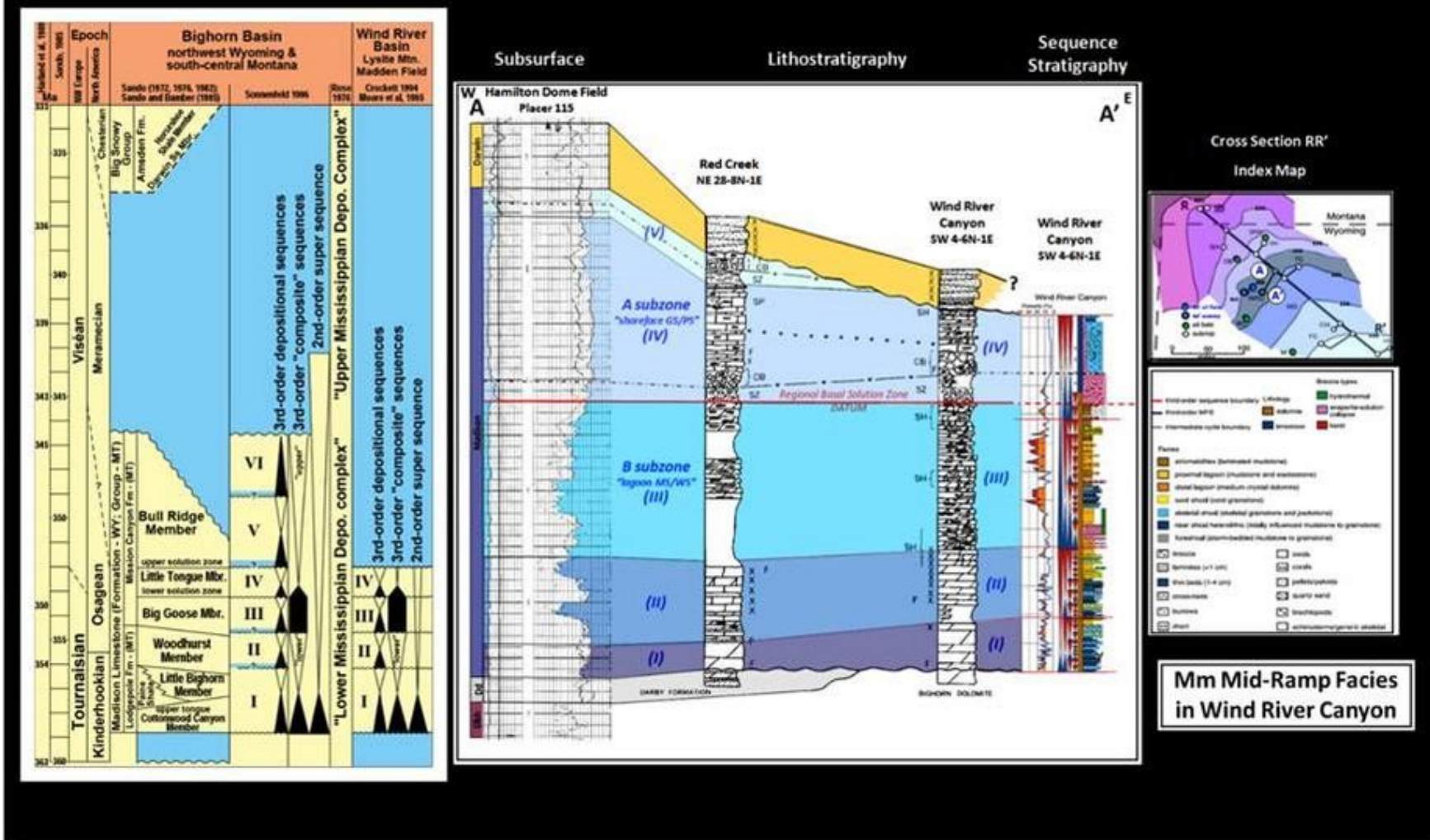
Left: Paleography of North America in the Early Mississippian. Right: Paleogeography of Mississippian Madison platform during deposition of the Little Tongue Member (Sequence IV) the uppermost unit present in Wind River Canyon.

Images: Left: After Blakey, R., 2011, NAM_key-345Ma_EMiss, Key Time Slices of North America Geologic History DVD; Right: After Sonnenfeld, M. D., 1996, SEPM, Fig. 1, p. 167 (after Gutschick, Sandberg and Sando, 1980), in Moore, C. H., 2001, Carbonate Reservoirs, Developments in Sedimentology 55: Elsevier, Fig. 10.1, p. 342 (from CD).



Regional time/lithology profile showing 19 million year duration of the Madison exposure in the Wind River Canyon (WRC).

*Image: After Eldam, N.S., 2012, Structural Controls on Evaporite Paleokarst Development: Mississippian Madison Formation, Bighorn Canyon Recreation Area, Wyoming and Montana Uplifted Area: M.S. Thesis, University of Texas at Austin, Fig. 1.4, p. 9;
<https://repositories.lib.utexas.edu/handle/2152/ETD-UT-2012-05-5134>
 (https://repositories.lib.utexas.edu/handle/2152/ETD-UT-2012-05-5134)*



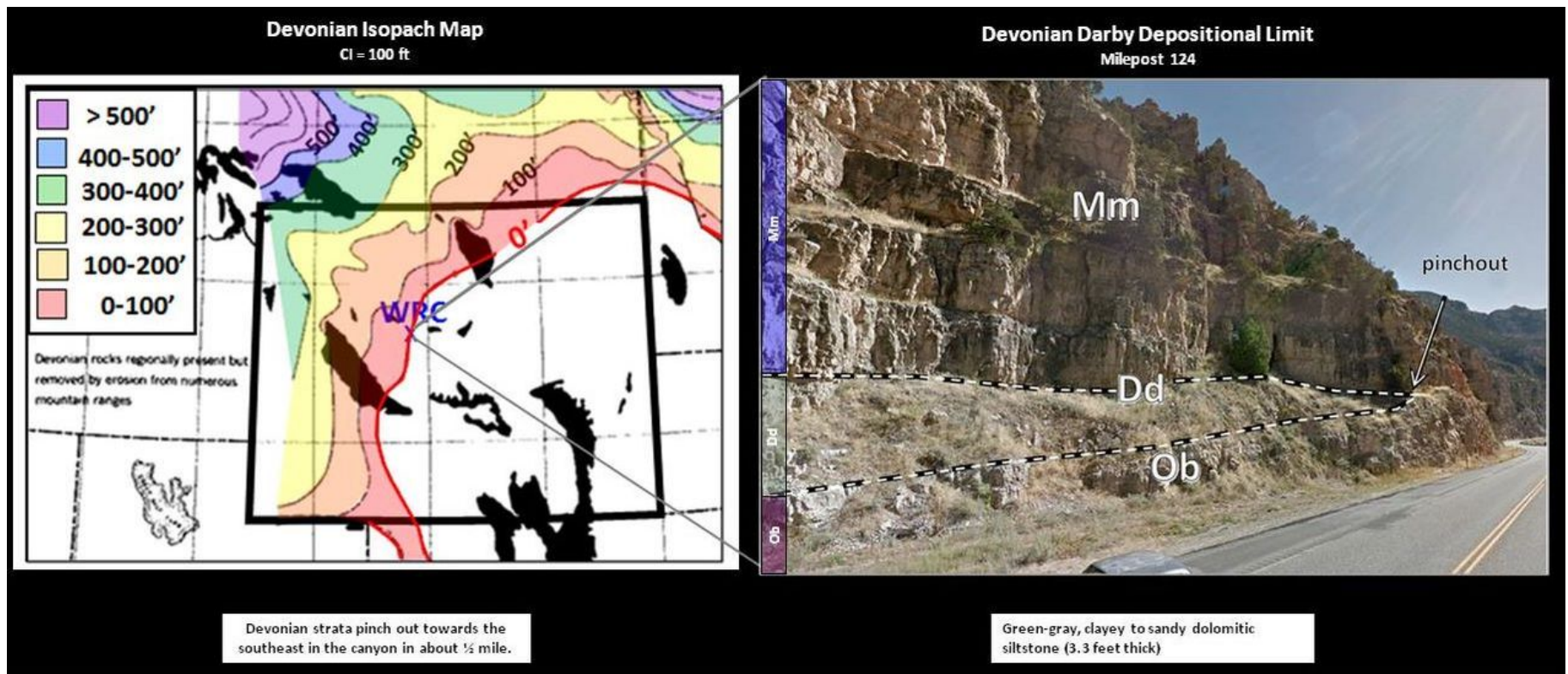
Madison Formation sequence stratigraphic column for the Bighorn/Wind River Basin. Right: Hamilton Dome subsurface log correlation to lithostratigraphic and sequence stratigraphic outcrop descriptions in southern Bighorn Basin. Cross section datum is the bottom of the regional solution zone at the base of Sequence IV (A Zone).

Images: Right: Moore, C. H., 2001, *Carbonate Reservoirs, Developments in Sedimentology* 55: Elsevier, Fig. 10.4, p. 345 (from CD); Right: Log: After Wyoming oil and Gas Conservation website <http://wogcc.state.wy.us/legacywogcce.cfm> (<http://wogcc.state.wy.us/legacywogcce.cfm>); Lithostratigraphic Columns: After Sando, W.J., 1967, *Madison Limestone (Mississippian), Wind River, Washakie and Owl Creek Mountains: AAPG Bulletin*, Vol. 51, no. 4, Fig. 4, p. 536; Sequence Stratigraphic Column: After Smith, T., and Eberli, G., 2004, *The Sequence Stratigraphic and Paleogeographic Distribution of Reservoir-Quality Dolomite, Madison Formation, Wyoming and Montana*, Fig. 1, p. 4; <http://mgi.rsmas.miami.edu/groups/csl/secure/archive/2001/Taury.pdf> (<http://mgi.rsmas.miami.edu/groups/csl/secure/archive/2001/Taury.pdf>)

The Madison Limestone post-depositional conversion to dolomite appears to depend on the original rock fabric. The finer grain rocks are preferentially dolomitized, which can increase the reservoir quality (porosity) of carbonates.

1.0 7.4 0.9 10.4

Top Ordovician Bighorn Dolomite. The Bighorn is approximately 140 feet thick in the Wind River Canyon. Some oil has been produced from the Bighorn in the Bighorn basin. At this location there is a three to four foot unit between the Madison and Bighorn called the Darby. This unit appears to be a filled channel and is composed of yellow slightly dolomitic mudstone and green shale. The Darby Formation produces at Garland, Hamilton Dome and Elk Basin. Milepost 124 marks the depositional limit of Devonian rock in central Wyoming.



Devonian Darby Formation depositional limit. Left: Thickness map shows 0-edge in Wind River Canyon (WRC). Right: Outcrop photo of formation pinchout. Mm is Mississippian Madison, Dd is Devonian Darby and Ob is Ordovician Bighorn

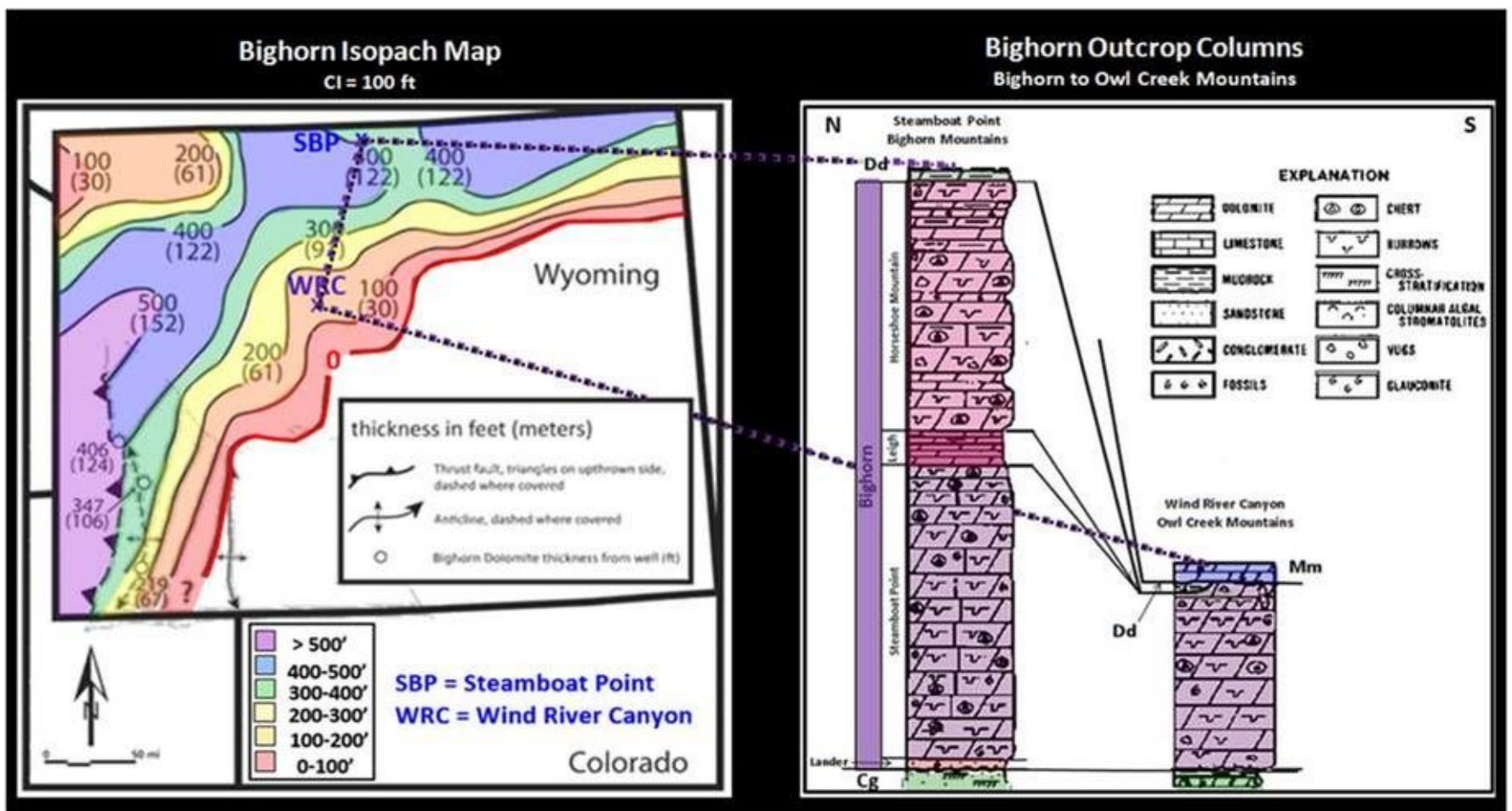
Images: Left: After Baars, D., 1972, RMAG Geologic Atlas, Fig. 4, p. 94. Right: After <http://usgeologymorphology.com/wind-river-43-32-45-N-SV-800px.jpg>://
<http://usgeologymorphology.com/wind-river-43-32-45-N-SV-800px.jpg>
<http://usgeologymorphology.com/wind-river-43-32-45-N-SV-800px.jpg>

The Middle to Late Ordovician Bighorn Dolomite forms the lowermost prominent carbonate cliffs in the Wind River Canyon. Thickness ranges from 85 to 140 feet. The formation was deposited in a regionally extensive tropical marine shelf environment that extended from Greenland to New Mexico. The lowest massive Steamboat Point Member is comprised of mottled light to medium gray and medium to coarse grained crystalline dolostone with locally

abundant chert nodules. Incomplete dolomitization results in variable weathering producing the characteristically hackly weathered surface perpendicular to the stratification on outcrop (Kilsgaard et al., 1972). Cyclic bedding on a scale of about 1–3 feet is visible from a distance and has been interpreted as the result of glacial-induced sea level fluctuations.

Chimney Rock near base of Bighorn exhibits the typical hackly texture of the dolomite. Top Cambrian Gallatin Formation just below Chimney Rock on treed slope as road veers eastward (L/R). The abrupt change in lithology from the Bighorn to the Gallatin is evident wherever it outcrops on south to Boysen Dam. "The Gallatin Formation is about 450 feet thick and is composed of gray and greenish shales at the base, with a middle cliff-forming limestone member (Pilgrim Limestone) about 80 feet thick, overlain by greenish shales (largely covered) and an upper varicolored reddish, purplish and greenish limestone, characterized by flat pebble limestone conglomerate intervals. These flat-pebble conglomerates are interpreted as rip-up clast deposits produced by storms that occasionally agitated the generally low energy, shallow water marine environments characteristic of this period.

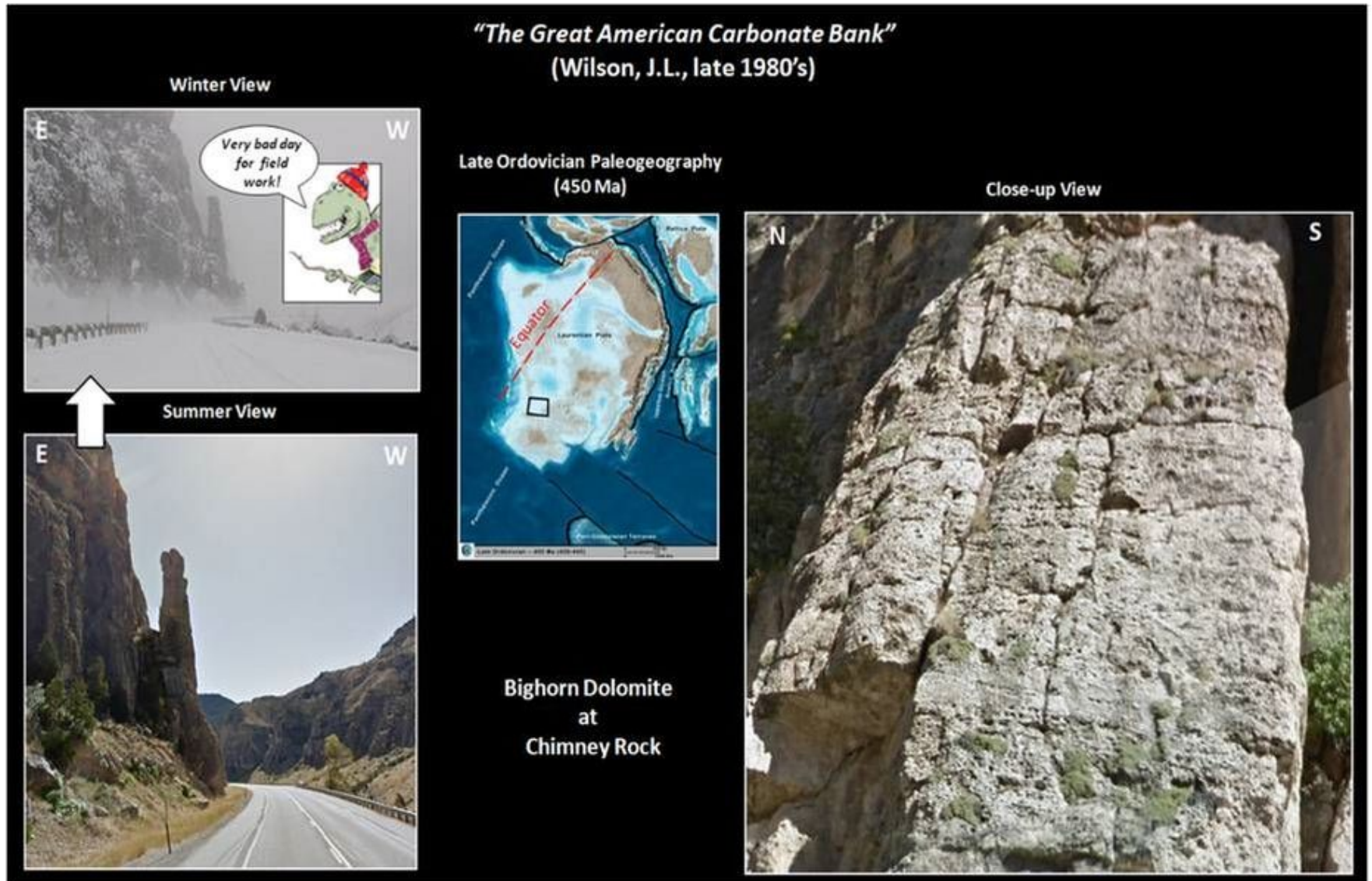
0.9 8.3 0.2 9.5



Ordovician Bighorn Formation. Left: Thickness map shows 0-edge about 25 miles south of Wind River Canyon (WRC). Right: Bighorn Dolomite outcrop cross section of from the Bighorn

to Owl Creek Mountains.

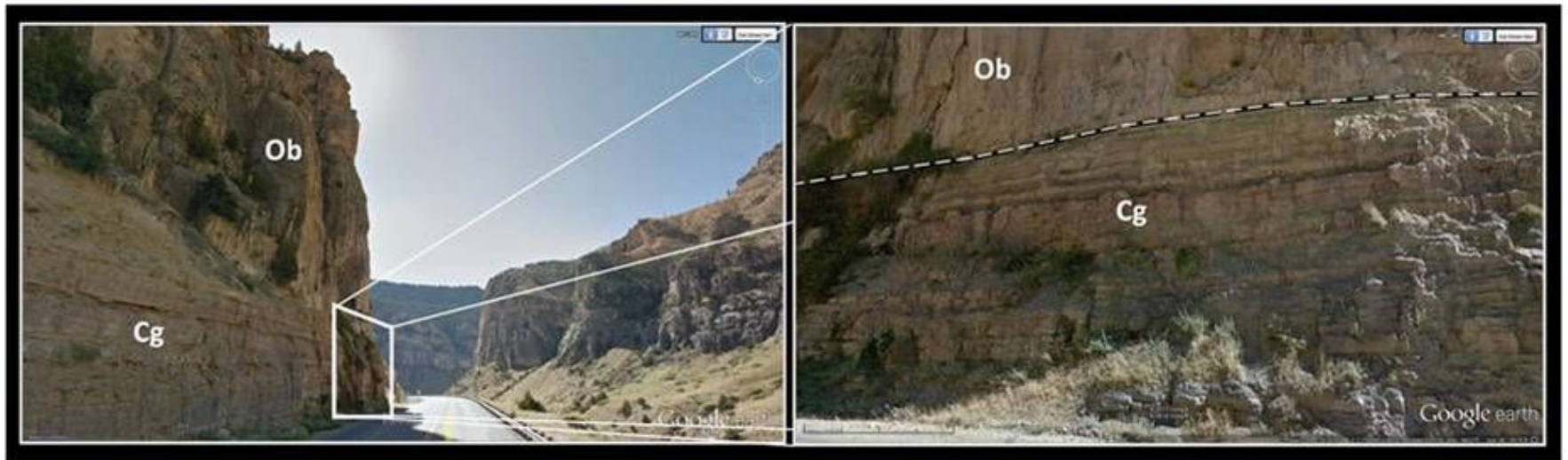
Images: Left: After Lynds, R. Campbell-Stone, E., Becker, T.P. and Frost, C.D., 2010, Stratigraphic evaluation of reservoir and seal in a natural CO₂ field: Lower Paleozoic, Moxa Arch, southwest Wyoming: *Rocky Mountain Geology*, v. 45, Fig. 4, p. 118; Right: After Zenger, D.H. and LeMone, D.V., 1995, Widespread "Bighorn facies", Upper Ordovician, North America; in Cooper, J.D., Droser, M.L., and Finney, S.C. (eds.), *Ordovician Odyssey: Short Papers for the Seventh International Symposium on the Ordovician System, Pacific Section, SEPM, Book 77, Fig. 3, p. 390.*



Ordovician Bighorn Dolomite, Chimney Rock, Wind River Canyon. Chimney Rock is at 8.3 miles.

Images: Left: Chimney Rock: Google Earth Street View in winter (top) and summer (bottom). Center: After Blakey, R., 2011, NAM_key-450Ma_LOrd, Key Time Slices of North America Geologic History DVD; Right: Chimney Rock close-up: Google Earth Street View

0.2 8.5 1.3 9.3 Good exposure of the Bighorn - Gallatin contact.

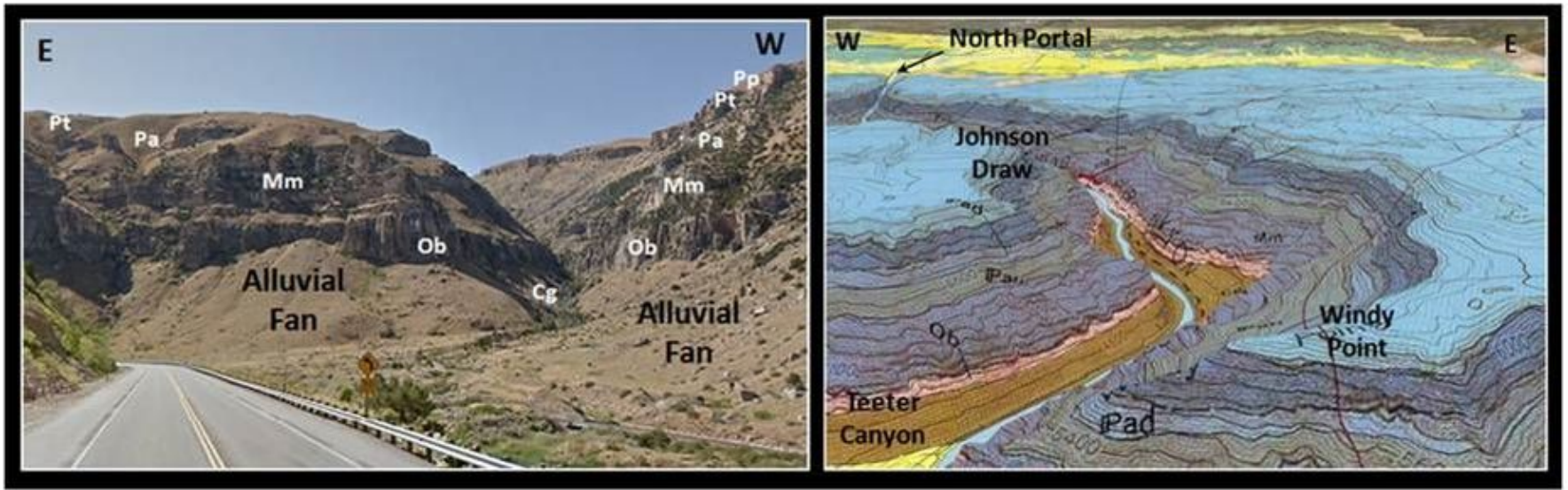


Ordovician Bighorn – Cambrian Gallatin contact on east wall and to southward in Canyon after sharp veer to the east (L/R) just past Teeter Canyon. Ob is Ordovician Bighorn and Cg is Cambrian Gallatin. View is at 8.5 miles.

Images: Google Earth Street View

1.3 9.8 1.1 8.0

Highway veers to the sharply to the east around Windy Point. Teeter Canyon to the left has alluvial fans at the foot of the gulch.



Left: Southwest view into Teeter Canyon. Right: Oblique north aerial view of geologic map near Teeter Canyon and Windy Point. Geologic notation: Pp: Permian Phosphoria Formation, Pt: Pennsylvanian Tensleep Formation, Pa: Pennsylvanian Amsden Formation, Pad: Ranchester Member of Pennsylvanian Amsden Formation, Pah: Horseshoe Member of Pennsylvanian Amsden Formation, Mm: Mississippian Madison Formation, Ob: Ordovician Bighorn Formation, Cg: Cambrian Gallatin Formation.

Images: Left: Google Earth Street View; Right: After Maughan, E.K., 1972, Geologic map of the Wedding of the Waters quadrangle, Hot Springs County, Wyoming: U.S. Geological Survey, Geologic Quadrangle Map GQ-1042, scale 1:24,000

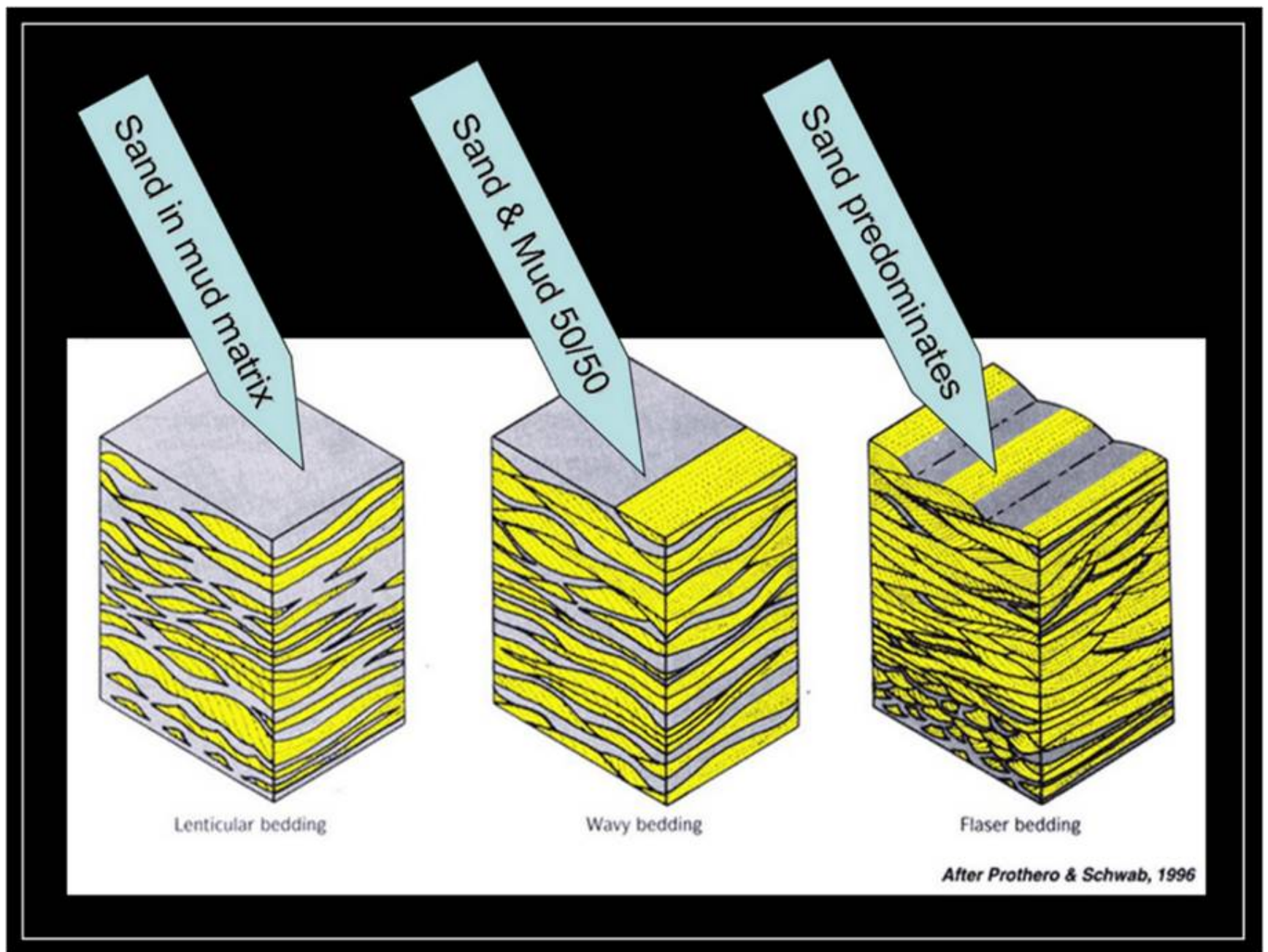
Approximate top of Cambrian Gros Ventre Formation (covered here). The Gros Ventre Formation is approximately 400 feet thick and consists mostly of medium to dark greenish gray micaceous and glauconitic mudstone. Clayey and silty sandstone beds occur in the lower portion of the formation, while calcareous mudstone and clayey limestone is commonly in the upper portion grading into the overlying Gallatin formation. Evidence for shallow water deposition of the Gros Ventre Formation includes ripples and horizontal burrows and tracks (many from trilobites) on bedding surfaces, and the presence of flaser bedding (bidirectional bedding pattern of alternating mud and sand, commonly associated with tidal currents).

1.1

10.9

1.3

6.9



Bidirectional, bedding patterns.

Image:

http://www.seddepseq.co.uk/SEDIMENTOLOGY/Sedimentology_Features/Strat_Bedforms/Bedding.htm#flaser

(http://www.seddepseq.co.uk/SEDIMENTOLOGY/Sedimentology_Features/Strat_Bedforms/Bedding.htm#flaser)

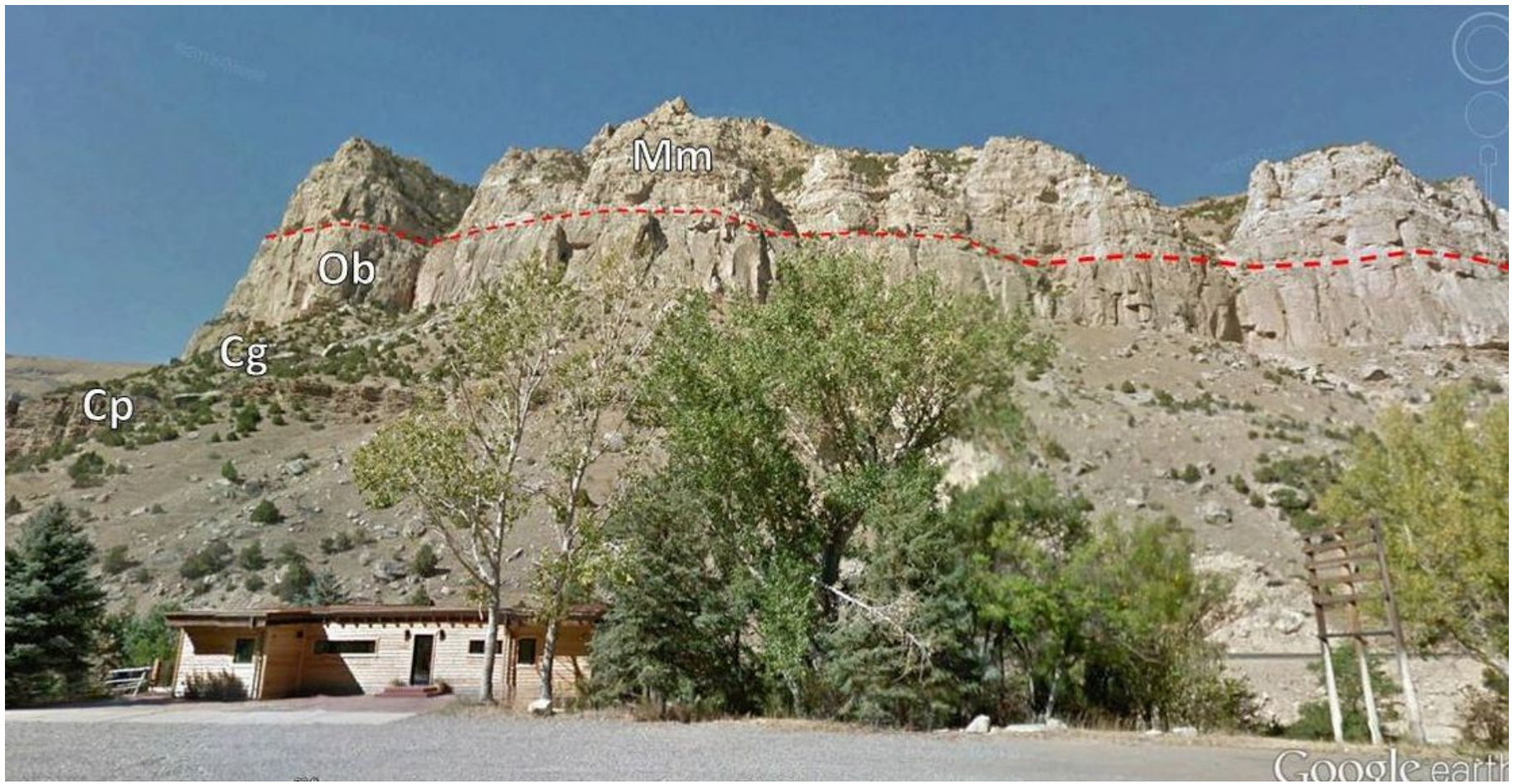
1.3

12.2

1.1

5.6

Old Wolf Creek Store, now private residence. The basal Gallatin Formation Pilgrim Limestone is well exposed on both sides of the canyon. Remains of numerous mudslides can be seen in this part of the canyon that damaged the railroad and highway during Spring 2015 when over six inches of rain fell in a week.



Madison (Mm) and Bighorn (Ob) on skyline on west side of highway. Slopes are Cambrian Gallatin (Cg) and Cambrian Pilgrim Limestone (Cp). This view is at 12.2 miles.

Image: Google Earth

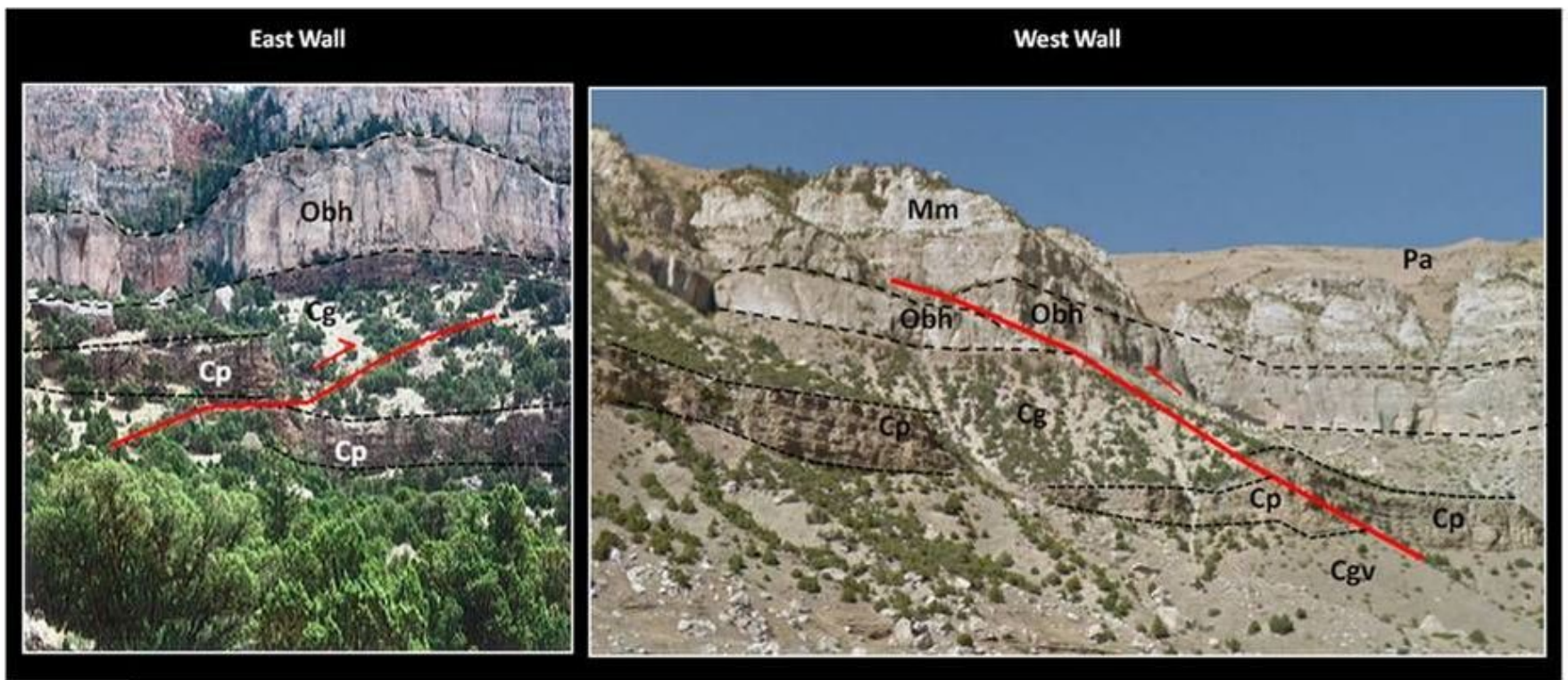
1.1

13.3

0.5

4.5

Second turnout to west (L/R).. Stop 4. Back limb thrust in lower Paleozoic rocks. Repeat section in the Pilgrim Limestone can be seen on both sides of the canyon.



Back-limb thrust in Wind River Canyon. Fault is the result of a bedding plane detachment within the incompetent Cambrian shales in response to structural folding during Laramide mountain building. Geologic notation: Pa: Pennsylvanian Amsden Formation, Mm: Mississippian Madison Limestone, Obh: Ordovician Bighorn Dolomite, Cg: Cambrian Gallatin Formation, Cp: Cambrian Pilgrim Limestone Member of Gallatin Formation.

Images: Steele, K.K., 2007, WRC Cp back-limb thrust fault.

0.5

13.8

0.5

4.0

Cambrian Gros Ventre - Flathead Sandstone contact. The Flathead is about 250 feet thick and composed mostly of sandstone and quartzite. Cross-bedding, worm burrows, and ripple marks can be found in the outcrop.



Contact of Cambrian Gros Ventre and Flathead Formations on east canyon wall.

Image: <http://www.wsgs.uwyo.edu/StratWeb/OwlCreekMts/OwlCreekMtsALL.aspx>
 (http://www.wsgs.uwyo.edu/StratWeb/OwlCreekMts/OwlCreekMtsALL.aspx)

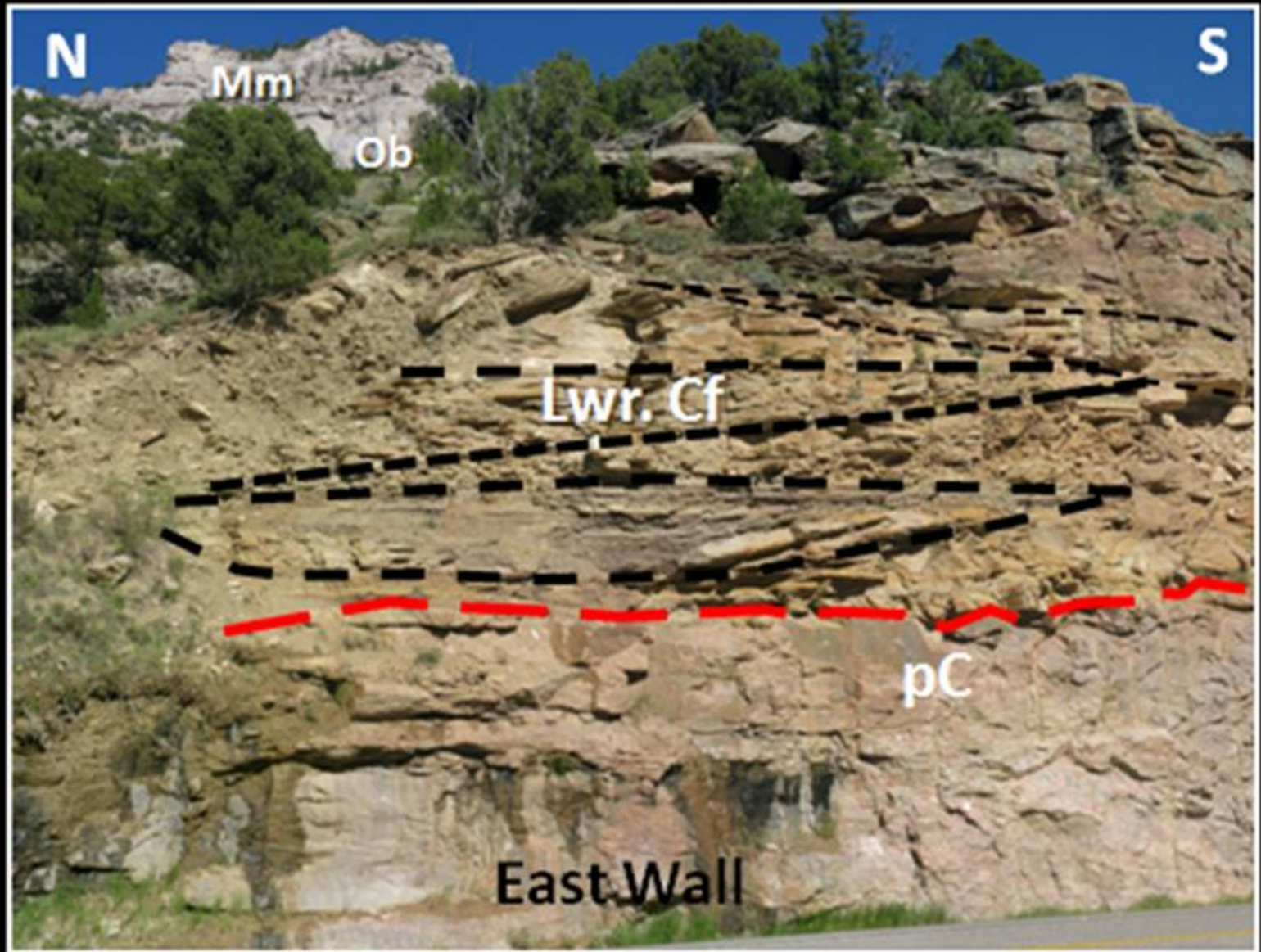
0.5

14.3

0.2

3.5

Top of Precambrian granite. Stop 5 (R/L). The weathered zone in top of the granite may be observed from here. Note channels in lower Flathead



Channels developed in the cross bedded Lower Cambrian Flathead overlying Precambrian granite. Contact is red dashed line. Geologic notation: Mm: Mississippian Madison Limestone, Ob: Ordovician Bighorn Dolomite, Lwr. Cf: lower member of Cambrian Flathead Sandstone.

Image: Bentley, C., 2011, Archean/Cambrian nonconformity at Wind River Canyon, Wyoming
<http://gigapan.com/gigapans/80213> (<http://gigapan.com/gigapans/80213>)

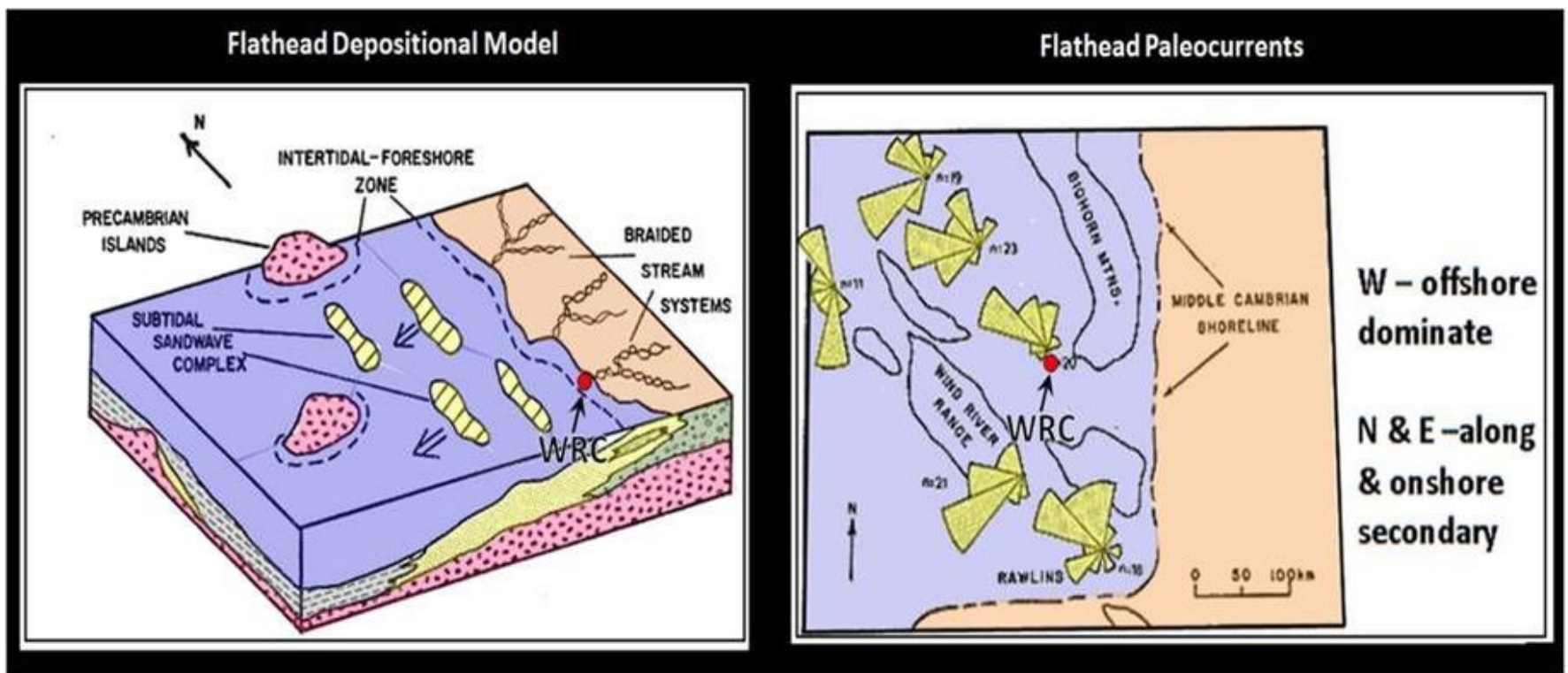
The Flathead Sandstone represents the first record of Phanerozoic sedimentation as the Middle Cambrian (520-505 million years ago) seas began to flood Wyoming. The shoreline is oriented north-south. The thickness of the Flathead reflects the topography on the Precambrian unconformity.



North American paleogeography and depositional facies in the Cambrian. At this time Greenland and part of Scotland were part of the continent. Location of Wind River Canyon (WRC) indicated by red dot. Tectonic feature notation: ADD: Adirondack Dome, CAT: Carolina Trench, CIA: Cincinnati Arch, ILB: Illinois Basin, MIB: Michigan Basin, OZD: Ozark Dome, WIA: Wisconsin Arch, WIB: Williston Basin

Images: Left: After Blakey, R., 2011, NAM_key-540Ma_ECam, Key Time Slices of North America Geologic History DVD; Right: After Ibid., cambrian-facies.

The Flathead is separated into a basal conglomerate and feldspathic sandstone deposited by braided streams, a deltaic flat-laminated feldspathic sandstone with shaley layers and a hummocky and cross-bedded feldspathic sandstone. The quartz and feldspar mineral components of the sands were sourced from Precambrian granitic bedrock.

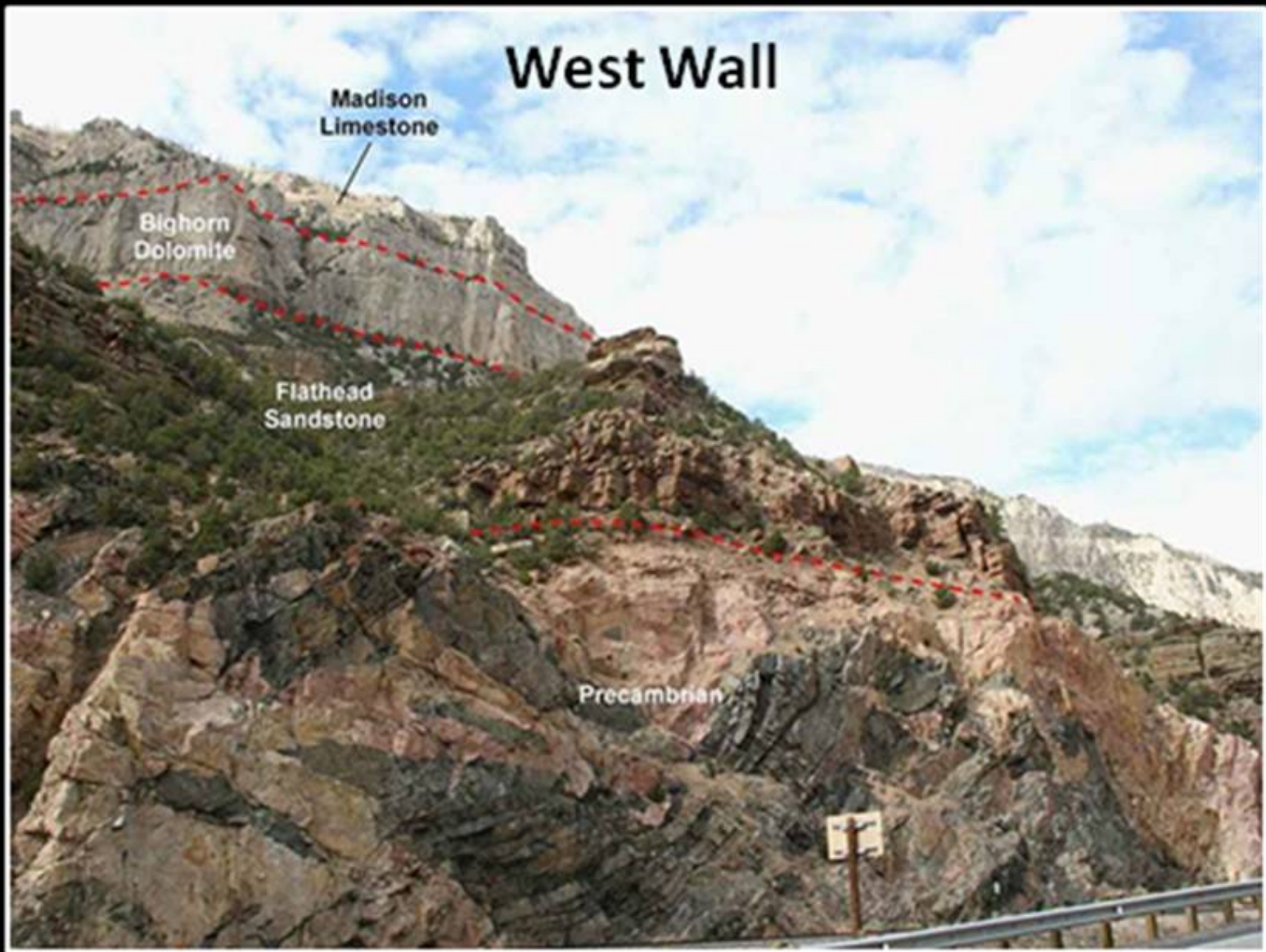


Flathead depositional model and paleocurrent data. Wind River Canyon (WRC) location shown by red dot.

Images: Left: After Middleton, L.D., 1980, *Sedimentology of middle Cambrian Flathead sandstone*, Wyoming PhD Dissertation, University of WY, Fig. 2.21, p. 71; Right: After Ibid., Fig. 2.22, p.72.

0.2 14.5 1.0 3.3

Precambrian complex of the Sweetwater subprovince. These Precambrian rocks represent the basement rock of Wyoming and are among the oldest rocks in North America. Darker rocks are black hornblende amphibolite and quartz biotite schist. Formed at about 2,900 Ma and metamorphosed around 2,750 Ma, this sequence was likely originally a series of medium to fine grained shale, siltstone, and sandstone and basalt and basaltic andesite flows and/or dikes and sills deposited on the margin of the ancient Wyoming Craton (Wyoming was a separate micro-continent at the time). Light colored rocks are pinkish gray to white quartz monzonite ("granite") intrusions some of which are pegmatites containing large crystals from cooling slowly. The dark schist and amphibolite make up the southern two thirds of the basement core of the Owl Creek Mountains.



Younger pink granite pegmatite intruding into Archean-age black schist and amphibolite, that were originally sediments and extrusive basalts deposited in the Sweetwater trough. This ancient trough developed at the margin of the Wyoming craton almost 3 billion years ago. It marks the suture between subprovinces within the craton.

Image:

<http://www.wsgs.uwyo.edu/research/stratigraphy/OwlCreekMts/OwlCreekMtsPALEO.aspx>
(<http://www.wsgs.uwyo.edu/research/stratigraphy/OwlCreekMts/OwlCreekMtsPALEO.aspx>)



Hand samples of the Wyoming craton.

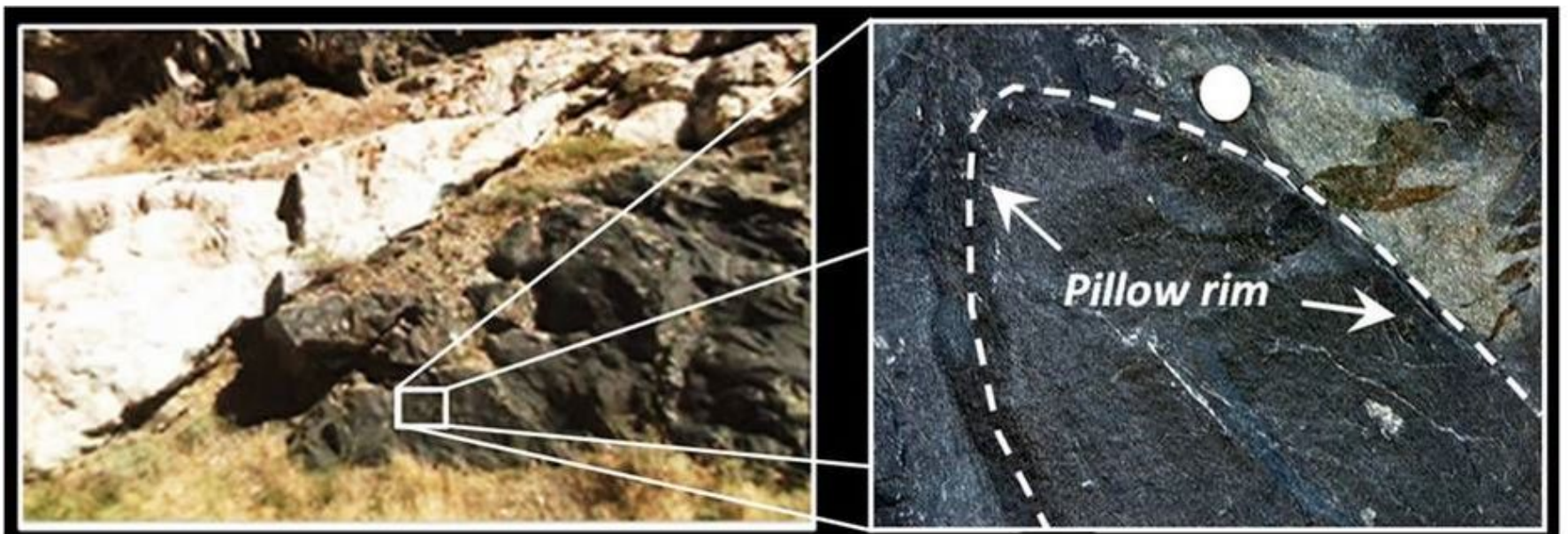
Images: Left: Granite, Wind River Canyon, Wyoming by Callan Bentley

<http://www.gigapan.com/gigapans/143006> (<http://www.gigapan.com/gigapans/143006>);

Right: Schist Ibid, <http://www.gigapan.com/gigapans/141715>

(<http://www.gigapan.com/gigapans/141715>)

1.0 15.5 0.1 2.3 2.9 Ga amphibolite - possible Archean pillow lava (Mid-Ocean Ridge Basalt) Protolith on east wall (L/R, before Tunnel #1).



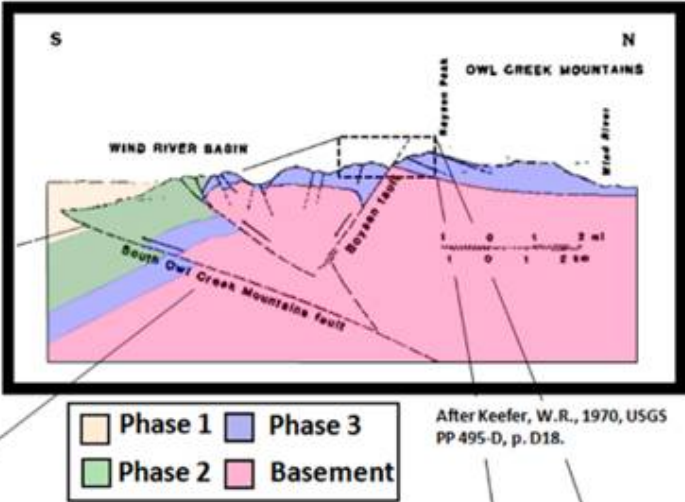
Precambrian remnant pillow structure in Archean amphibolite. Site is about 250 feet of tunnel number 1. The pillow structure indicates the original lava was deposited in water.

Images: Left: Google Earth Street View;
Right:
http://www.earth.edu.waseda.ac.jp/photogeology/200504rocky/WindrivCanyon/1K_DSC_2404.jpg
(http://www.earth.edu.waseda.ac.jp/photogeology/200504rocky/WindrivCanyon/1K_DSC_2404.jpg)

0.1	15.6	0.1	2.2	Tunnel #2
0.1	15.7	0.1	2.1	Tunnel #3
0.1	15.8	1.5	2.0	Stop 6 at Boysen State Park pull out to view Boysen Fault

Boysen Fault

An intricate system of normal faults is present along the south margin of the Owl Creek Mountains in the vicinity of Wind River Canyon (Fanshawe, 1939; Tourtelot and Thompson, 1948; Wise, 1963). These faults are apparently confined to the hanging-wall block of the South Owl Creek Mountains fault; most are aligned parallel to the mountain front, but some are transverse to this trend. There is no consistent pattern as to which sides were upthrown or downthrown, but in general, there appears to be a lowering of the marginal part of the range with respect to its crest (Wise, 1963). The most extensive fault is the Boysen fault, which trends nearly east across the canyon from a point 3 miles west of the channel of Wind River to a point 7 miles east. The down dropped block is to the south, and maximum stratigraphic displacement is about 2,000 feet. The normal faults are best explained as having resulted from the tensional fracturing of rigid rocks across the crest of the range as it was being uplifted and arched during Laramide deformation. Individual blocks then collapsed along the steep margin of the range, either at the same time or somewhat later than the arching (last 10 Ma).

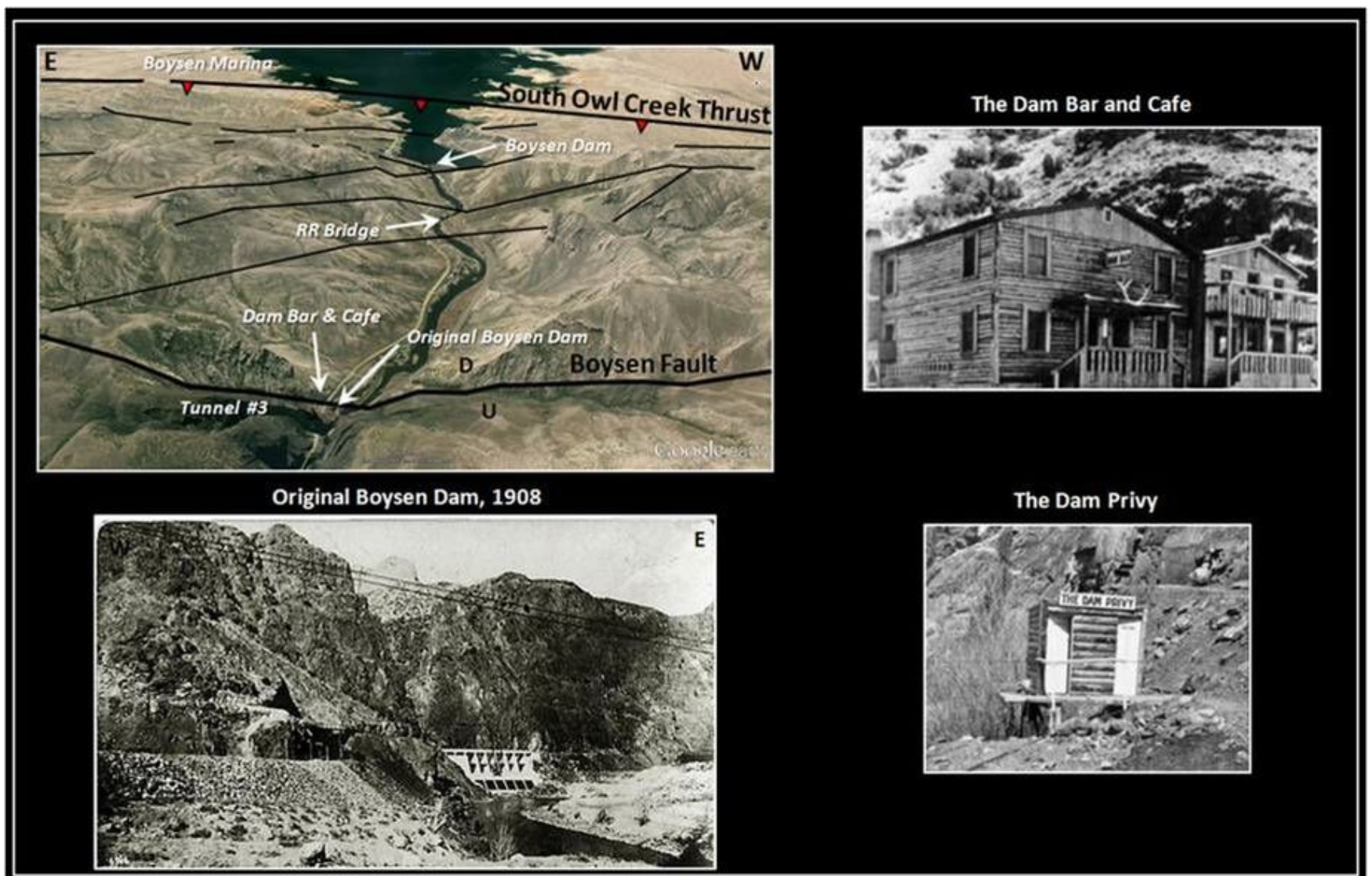


West view of the Boysen Normal Fault from Lower Wind River Campground. Boysen residence was located in this area.

Images: Cross section: After Keefer, W.R., 1970, USGS PP 495-D, p. D18; Bottom: After Mapcarta Street View.

Boysen Dam

The original Boysen Dam was located across the canyon on the upthrown side of Boysen Fault. The dam was constructed in 1908. Asmus Boysen envisioned a source of water and power for nearby mining operations. But events did not work out for his dam. Electricity was delivered to the mines, Shoshoni and Riverton beginning in 1911, but the railroad had begun legal proceedings against the power company in 1908. The 1915 Wyoming Supreme Court ruling required removal of the dam's super structure and made power generation unprofitable. The 1923 flood silted the plant generations and that ended its short existence. Mining activity slowed, more lawsuits arose and flooding continued to plague the railroad line in the canyon. After numerous legal battles the dam was finally removed in 1948 as part of constructing the new dam. Some of the abutments from the old dam can still be seen on the canyon walls on the south side of tunnel #3. The land upon which the upper and lower campgrounds are located was created by siltation from flooding. The new dam was completed in 1953.



Top Right: South aerial view of Boysen Dam area. Bottom Right View of original Boysen Dam,

1908. Top Right: Tourist facilities on the east side of highway just after Tunnel #3. Bottom Right: Old comfort station located near the bar and café.

Images: Upper Left: Google Earth. Dam and facilities: After Wilson, D., 2010, Boysen History Presentation <http://wyoparks.state.wy.us/pdf/HistoryofBoysenDam.pdf> (<http://wyoparks.state.wy.us/pdf/HistoryofBoysenDam.pdf>)

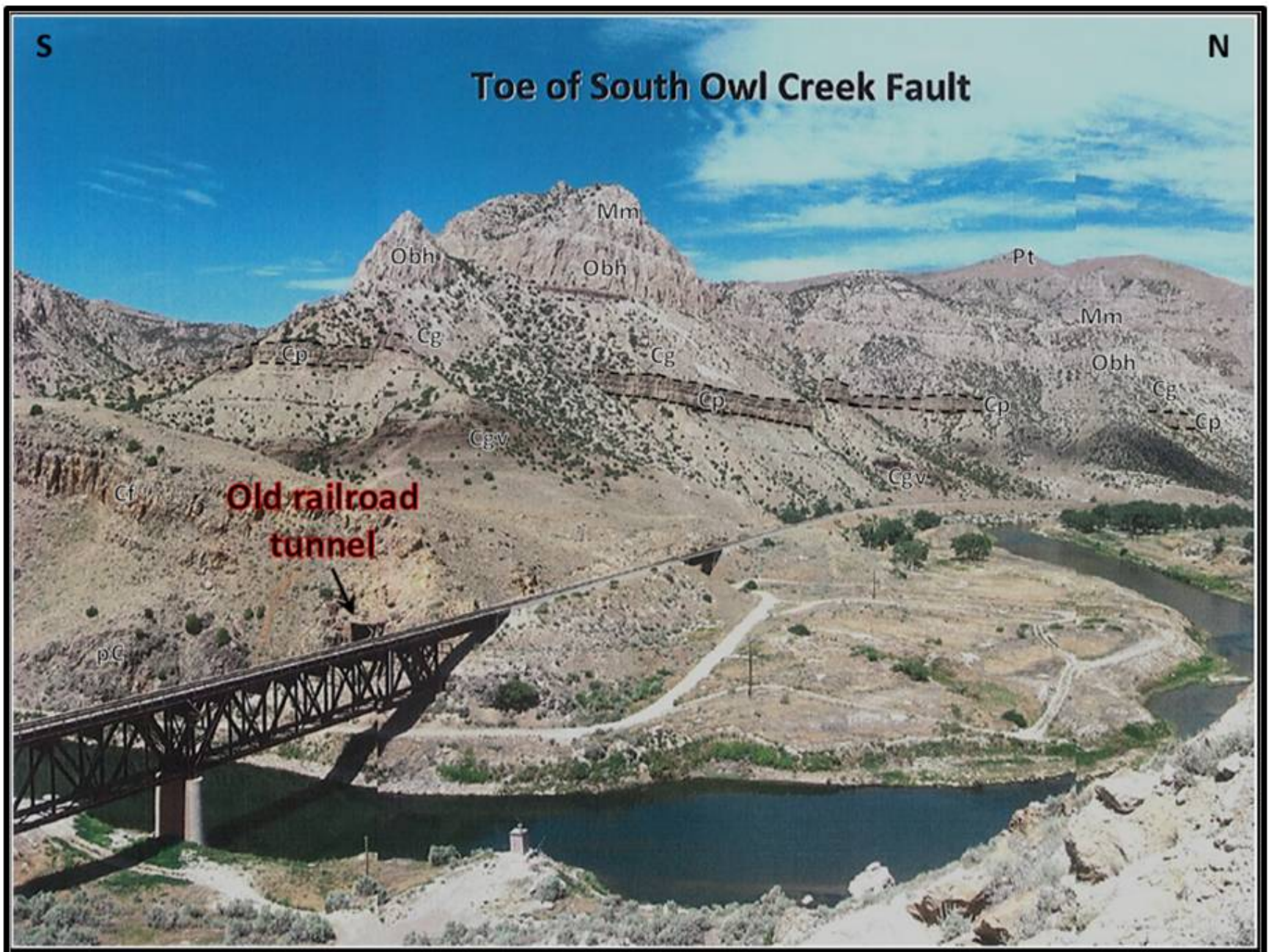
1.5

17.3

0.5

0.5

Stop 7 at gravel pad on right or west side of road. Walk to top of hill. The Boysen Structural complex is visible from this point.



Boysen structural complex. Numerous faults can be seen on this collapsed portion of the South Owl creek Thrust

Image: After Granath, J.W., 1975, Precambrian Research, Fig. 2, p. 3.

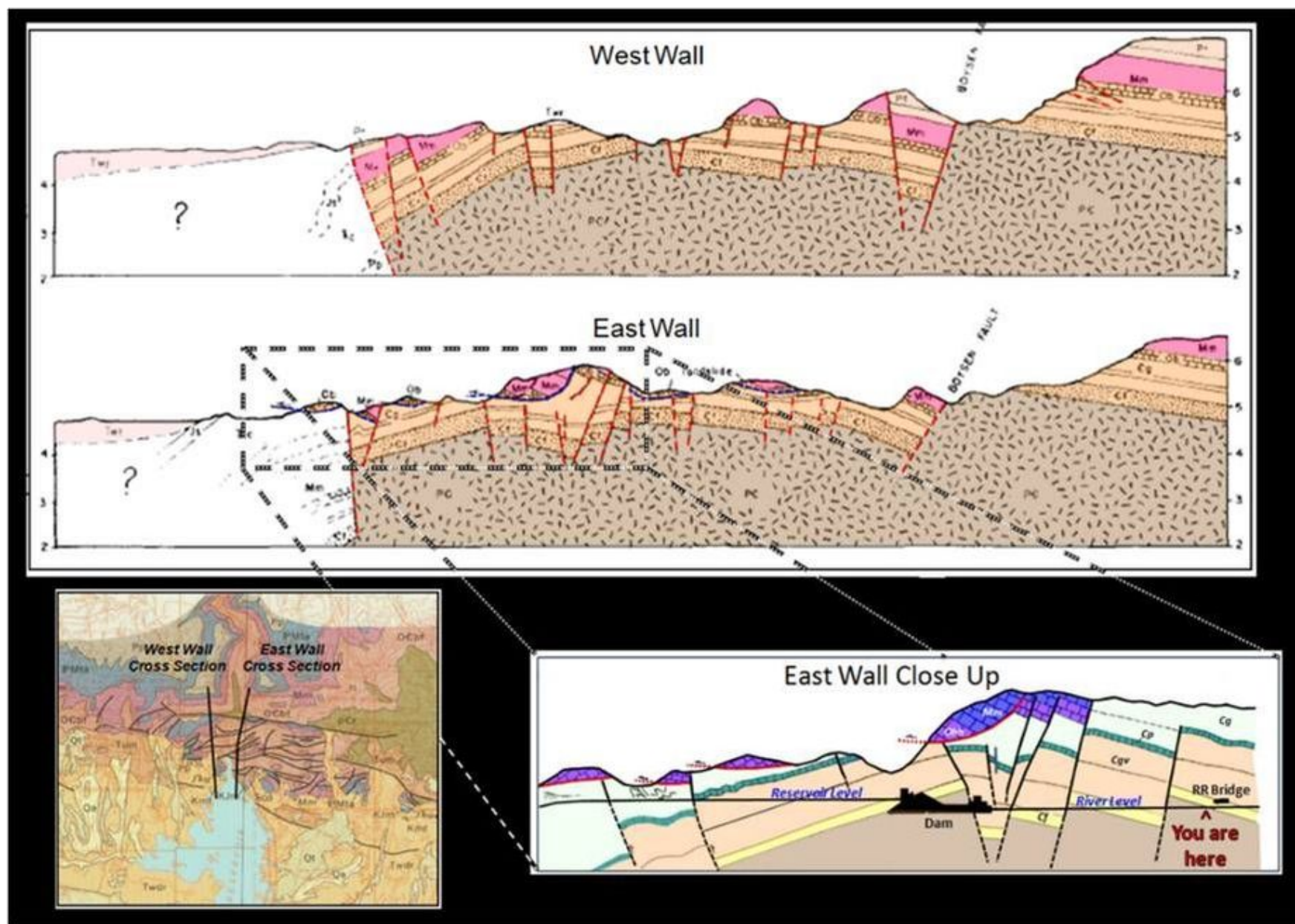
0.5

17.8

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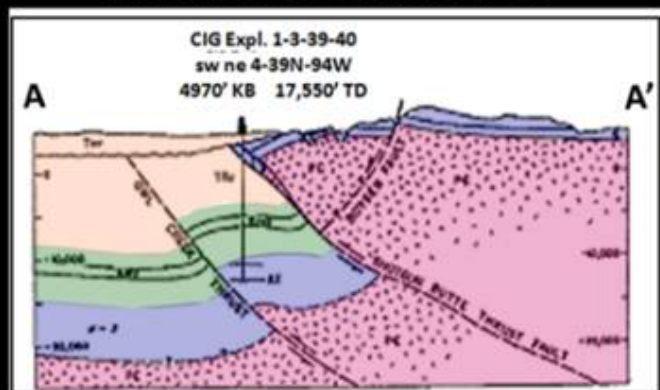
New Boysen Dam.



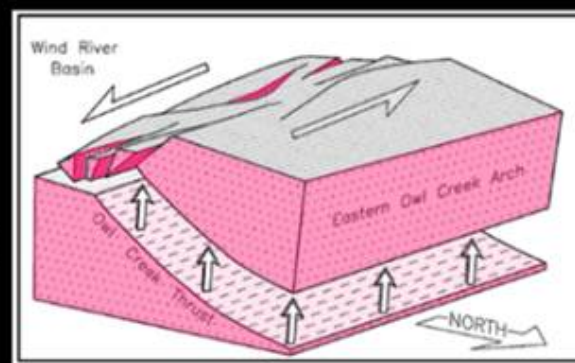
Cross sections in area of Boysen structural complex.

Images: After Wise, D.U., 1963, AAPG Bulletin Vol. 47, No. 4, Fig. 4, p. 592; Bottom Right: After Lageson, D. and Spearing, D., 1988, Roadside Geology of Wyoming, 2nd Ed.

Structural Cross Section of EOCA



Sinistral Oblique-slip Thrusting Model of EOCA



- 5 mile (26,000') Vertical Separation (Blackstone, 1990)
- 10 mile (52,000') Basement Overhang,
- 5.7 mile (> 30,000'); Structural Relief (Gries, 1983; Paylor and Lang, 1990);
- 1.25 mile (6,500+') Horizontal Separation (Caldwell, 1977)

South View of Southwestern BHB



Structure of the East Owl Creek Arch (EOCA) and the South Owl Creek Fault.

Images: Upper Left: After Blackstone, D.L., Jr., 1990, Rocky Mountain Foreland Structure Exemplified by the Owl Creek Mountains, Bridger Range and Casper Arch, Central Wyoming: WGA, Cross Section 14; Upper Left: After Molzer, P.C. and Erslev, E.A., 1995, Oblique Convergence During Northeast-Southwest Laramide Compression Along the East-West Owl Creek and Casper Mountain Arches, Central Wyoming: AAPG Bulletin, Vol. 79, No. 9, Fig. 16, p. 1393; Bottom: Google Earth.

End of Road Log



North aerial view of Boysen Reservoir and the south portal of the Wind River Canyon “Crack in Time.”

Image: Computer Terrain Mapping, Inc., 2013, Superimposed Rivers

http://www.ctmap.com/ctm_image/boysen.png

(http://www.ctmap.com/ctm_image/boysen.png)

Things To Do In Wind River Canyon & Thermopolis Area

The Wind River Canyon is on the Wind River Indian Reservation, so public access in the Canyon is restricted to the pullouts next to the highway. Our recommendations for fun things to do include fishing & rafting, and visits to Anchor Dam (<http://www.geowyo.com/anchor-dam.html>), Dinosaur Marbles (<http://www.geowyo.com/dinosaur-marbles.html>), Thermopolis Hot Springs (<http://www.geowyo.com/thermopolis-hot-springs.html>) and the Wyoming Dinosaur Center (<http://www.wyodino.org/>).

Fishing & Rafting – You can fish the river in the Canyon from the shore with a Wind River Tribal Fishing/Trespass Permit and the annual Recreation Stamp which you can buy in Thermopolis. Another option is guided whitewater rafting and fishing trips in the Canyon run by Wind River

Canyon Whitewater & Fly Fishing (<http://www.windrivercanyon.com/>) also located in Thermopolis. Just north of the Canyon and north of the Reservation, starting at the Wedding of the Waters public access, you can float and/or fish the Big Horn River. From Wedding of the Waters to Thermopolis, the river is flanked by private land, thus you will need a boat/raft or a guided trip to access this stretch. Inquire in town about river takeout options, river flows, and hazards which include low bridge clearance. Fishing this section requires a Wyoming Game & Fish (<https://wgfd.wyo.gov/fishing-and-boating>) license which you can also buy in Thermopolis.

Wind River Canyon

Whitewater Rafting



<http://www.windrivercanyon.com/>

Trout Fishing



<http://lh5.ggpht.com/-hVhOX52K2i0/VEMX4eII9nI/AAAAAAQvA/LOs51EgG9jU/s1600-h/Thermopolis1-7-Oct-2014b6.jpg>

Browns



<http://www.windrivercanyon.com/>

Rainbows



<https://www.yellowdogflyfishing.com/sites/default/files/styles/colorbox/public/trips/%C2%AC%20Klug%20Photos%20-%20Wind%20River%202015-30.jpg?itok=WMCA0K7Y>



<https://i.ytimg.com/vi/dxU5PmGO44U/maxresdefault.jpg>



http://riverdrifters.net/wp-content/uploads/photo-gallery/gallery_0002_wind.jpg

Hot Springs State Park (<http://wyoparks.state.wy.us/index.php/places-to-go/hot-springs>) – Within the state park there are trails to the river, springs, terraces, and overlooks. There is the State Bath House and two hot mineral water commercial swimming facilities at Hellie’s Tepee Pools and Star Plunge (<http://www.starplunge.com/>).



http://farm8.staticflickr.com/7183/6803600488_9bd71ae04b.jpg



<https://tornadogirl.files.wordpress.com/2010/06/storm-chasing-may-22-2010-sd-018.jpg>



http://wyoparks.state.wy.us/sphst/Parks_Sites_Trails/images/com_droppics/436/Hot-Springs-087.jpg

Hot Springs State Park Thermopolis WY



https://c1.staticflickr.com/7/6124/6043890487_43fcc26fd3_b.jpg



http://img1.sunset.timeinc.net/sites/default/files/styles/A_3_horizontal_inbody_900x506/public/image/2016/09/main/thermopolis-wyoming-star-plunge-pools-0811.jpg?itok=DdUe2ftY



http://res.cloudinary.com/miles-extranet-dev/image/upload/ar_16:9,c_fill,w_1200,g_face/Wyoming/account_photos/1096/f5ab66db73a7d551a4908ed0ab10ab30_HSSP-bathhouse.jpg



<https://www.thermopolisdaysinn.com/images/5.png>

Wyoming Dinosaur Center (<http://www.wyodino.org/>) – In addition to dinosaur exhibits and the bone preparation lab, they have a fine museum with a wide variety of fossils over geologic history displayed in a “Walk Thru Time”. From late-May through mid-September, they offer tours of the dig site and a “Dig For A Day” program where you can help them dig in the dinosaur bone beds.

Wyoming Dinosaur Center Thermopolis WY



<http://dinosaurstop.com/wp-content/sabai/File/files/3a7cccd5071518813f23002278a8b.jpg>



<https://media-cdn.tripadvisor.com/media/photo-s/01/2d/5f/7f/doesn-t-look-like-much.jpg>



<https://media-cdn.tripadvisor.com/media/photo-s/06/7f/89/30/wyoming-dinosaur-museum.jpg>



<http://www.wyodino.org/wp-content/uploads/2015/12/Exploring-1.jpg>



<http://dinosaurstop.com/wp-content/sabai/File/files/3d1b86ff7a1190b32f260795a25370cb.jpg>

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